

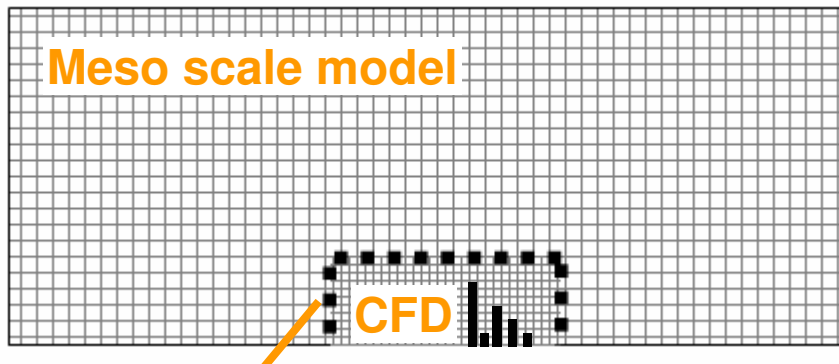
Adaptation of Pressure Based CFD Solvers for Mesoscale Atmospheric Problems

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4-th May 2009.



Advantages of a CFD based model

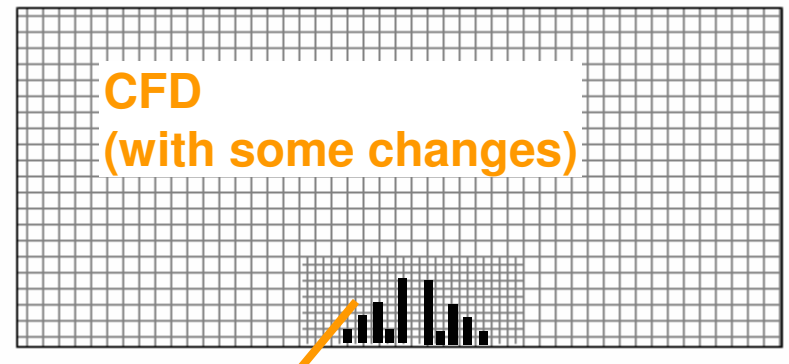
Meso scale model



model conversion interface

- The bidirectional interface is a source of numerical errors eg. it can cause partial reflection.

CFD
(with some changes)



grid refinement

- Better geometrical description
- More general turbulence models
- Easy customization
- Advanced pre- and post processing

Gravity waves ??

Thermal convection (UHIC) ??



Methodology

Incompressible CFD model (FLUENT)
+ transformation system
+ customized source terms



Mathematical description

$$\bar{\rho} = \rho_0 - \rho_0 \beta (\tilde{T} - T_0)$$

$$\nabla \cdot \bar{\mathbf{v}} = 0$$

$$\frac{\partial}{\partial t} (\rho_0 \bar{\mathbf{v}}) + \nabla \cdot (\rho_0 \bar{\mathbf{v}} \otimes \bar{\mathbf{v}}) = -\nabla \bar{p} + \nabla \cdot \boldsymbol{\tau} + (\bar{\rho} - \rho_0) \mathbf{g} + \mathbf{F}$$

Customized
volume sources

$$\frac{\partial}{\partial t} (\rho_0 c_p \tilde{T}) + \nabla \cdot (\bar{\mathbf{v}} \rho_0 c_p \tilde{T}) = \nabla \cdot (K_t \nabla \tilde{T}) + S_T$$

$$\frac{\partial}{\partial t} (\rho_0 k) + \nabla \cdot (\rho_0 \bar{\mathbf{v}} k) = \nabla \cdot \left(\frac{\mu_t}{\sigma_k} \nabla k \right) + G_k + G_b - \rho_0 \varepsilon + S_k$$

$$\frac{\partial}{\partial t} (\rho_0 \varepsilon) + \nabla \cdot (\rho_0 \bar{\mathbf{v}} \varepsilon) = \nabla \cdot \left(\frac{\mu_t}{\sigma_\varepsilon} \nabla \varepsilon \right) + \rho_0 C_{1\varepsilon} S \varepsilon - \rho_0 C_2 \frac{\varepsilon^2}{k + \sqrt{\nu \varepsilon}} +$$

$$+ C_{1\varepsilon} \frac{\varepsilon}{k} C_{3\varepsilon} G_b + S_\varepsilon$$

Transformed variables

$\bar{\rho}, \tilde{T}, \bar{p}, \bar{\mathbf{v}}, \mathbf{z}$

Transformation expressions

$$T = \tilde{T} - T_0 + \bar{T}$$

$$p = \frac{\bar{\rho}}{\rho_0} \cdot \tilde{p} + \bar{p} = e^{-\zeta z} \cdot \tilde{p} + \bar{p}$$

$$\rho = \tilde{\rho} - \rho_0 + \bar{\rho}$$

$$z = -\frac{1}{\zeta} \text{Ln}(1 - \zeta \tilde{z})$$

$$w = \frac{\rho_0}{\bar{\rho}} \tilde{w} = \tilde{w} e^{\zeta z}$$

Equilibrium profiles

for proper elimination of the hydrostatic pressure gradients

$$\bar{T} = T_0 - \gamma z$$

$$\bar{p} = p_0 \left(\frac{T_0 - \gamma z}{T_0} \right)^{\frac{g}{R\gamma}}$$

$$\bar{\rho} = \rho_0 e^{-\zeta z}$$

$$T_0 = 288.15 \text{ K}$$

$$p_0 = 1.01325 \cdot 10^5 \text{ Pa}$$

$$\rho_0 = 1.225 \text{ kg/m}^3$$

$$\gamma = 0.65^\circ\text{C}/100\text{m}$$

$$g/(R\gamma) = 5.2553$$

$$\zeta = 10^{-4} \text{ m}^{-1}$$

Standard ISA profile (up to 11km)

Approximate profile

Error bound is within
0.4% below 4000 m.



Summary of source terms

In momentum equation:

$$S_u = \rho_0 f v - \rho_0 \ell \tilde{w} J$$

$$S_v = -\rho_0 f u$$

$$S_w = \rho_0 (J^2 - 1) (\ell u J^{-1} + \beta (\tilde{T} - T_0) g) + \rho_0 \ell u J^{-1} + \zeta J (\tilde{p} - \rho_0 \tilde{w}^2)$$

In the energy equation:

$$S_T = J S_\Theta - \rho_0 c_p \tilde{w} (\Gamma - \gamma) J$$

In the transport equation of turbulent kinetic energy

$$S_k = -\beta g \frac{\mu_t}{Pr_t} (\Gamma - \gamma)$$

In turbulent dissipation equation:

$$S_\varepsilon = -C_{1\varepsilon} C_{3\varepsilon} \frac{\varepsilon}{k} \beta g \frac{\mu_t}{Pr_t} (\Gamma - \gamma)$$

Stratification
+ adiabatic heating



$$\Gamma - \gamma = 0.33 \frac{^\circ\text{C}}{100\text{m}}$$



$$\ell = 2 \Omega \cos \phi$$



$$f = 2 \Omega \sin \phi$$



$$J = (1 - \zeta z)^{-1}$$

Coriolis force

Compressibility



Related publications

- [1] **Kristóf G, Rácz N, Balogh M: Adaptation of Pressure Based CFD Solvers for Mesoscale Atmospheric Problems, *Boundary-Layer Meteorol*, 2008.**
- [2] **N.Rácz, G.Kristóf, T.Weidinger, M.Balogh:** Simulation of gravity waves and model validation to laboratory experiments, *CD, Urban Air Quality Conf. Cyprus, 2007.*
- [3] **G.Kristóf, N.Rácz, M.Balogh:** Adaptation of pressure based CFD solvers to urban heat island convection problems, *CD, Urban Air Quality Conf. Cyprus, 2007.*
- [4] **G.Kristóf, N.Rácz, Tamás Bányai, Norbert Rácz:** Development of computational model for urban heat island convection using general purpose CFD solver, *ICUC6, 6-th Int.Conf.on Urban Climate, Göteborg, pp. 822-825., 2006.*
- [5] **G. Kristóf, T. Weidinger, T. Bányai, N. Rácz, T.Gál, J.Unger:** A városi hősziget által generált konvekció modellezése általános célú áramlástan szoftverrel - példaként egy szegedi alkalmazással, *III. Magyar Földrajzi Konferencia, Budapest, 2006., Bp, CD*
- [6] **Kristóf G., Rácz N., Bányai T., Gál T., Unger J., Weidinger T.:** A városi hősziget által generált konvekció modellezése általános célú áramlástan szoftverrel– összehasonlítás kisminta kísérletekkel *A 32. Meteorológiai Tudományos Napok előadásai. Országos Meteorológiai Szolgálat, Bp., 2006*
- [7] **Dr. Lajos T., Dr. Kristóf G., Dr. Goricsán I., Rácz N.:** Városklíma vizsgálatok a BME Áramlástan Tanszékén, hősziget numerikus szimulációja *VAHAVA projekt (A globális klímaváltozás: hazai hatások és válaszok) zárókonferenciája Bp. CD, 2006*
- [8] **Rácz N. és Kristóf G.:** Hősziget cirkuláció kisminta méréseinek összehasonlítása saját fejlesztésű LES modellel *Egyetemi Meteorológiai Füzetek No. 20 ELTE Meteorológiai Tanszék, Bp. 173-176, 2006.*
- [9] **M. Balogh, G. Kristóf:** Automated Grid Generation for Atmospheric Dispersion Simulations, *pp.1-6., MICROCAD konferencia, Miskolc, 2007.*

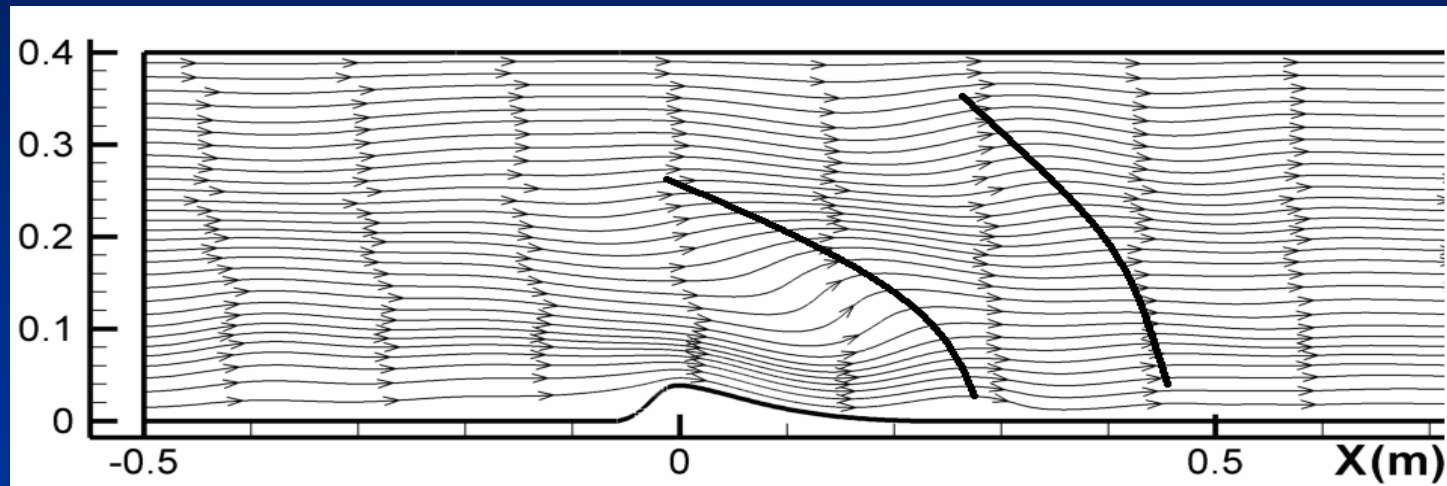
Model validation

1. analytical solutions
2. laboratory experiments
3. a standard test case
4. a full scale event

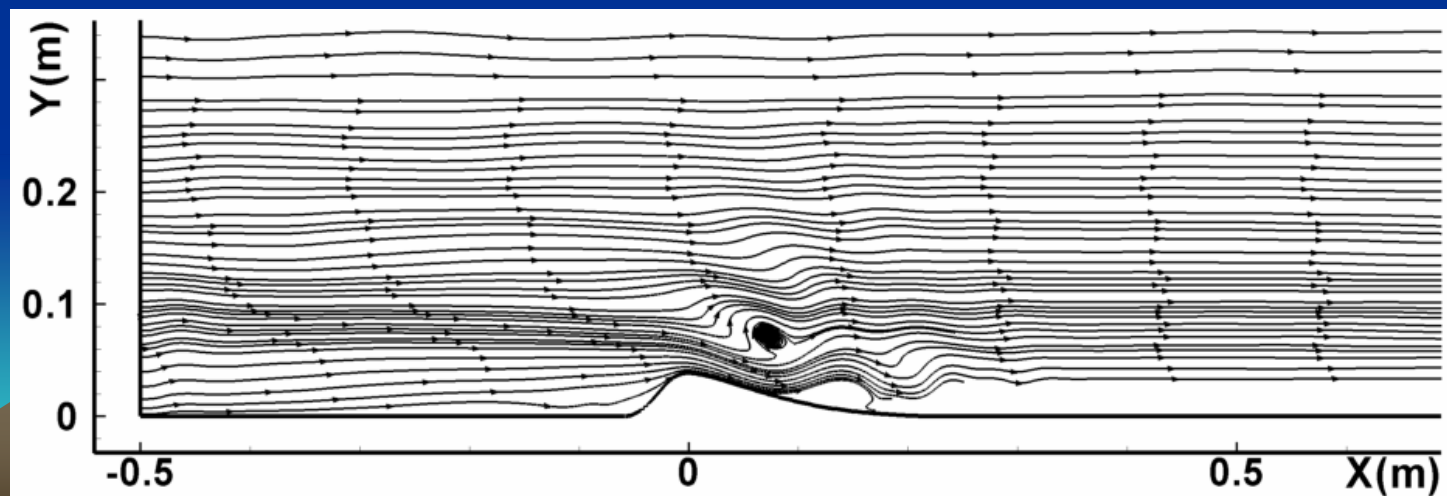
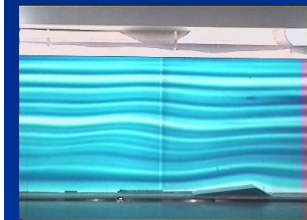


Gravity waves

Gyüre, B. and János, I.M., 2003. Stratified flow over asymmetric and double bell-shaped obstacles. *Dynamics of Atmospheres and Oceans* 37, 155-170.



$U/Nh = 1.4$



$U/Nh = 0.3$

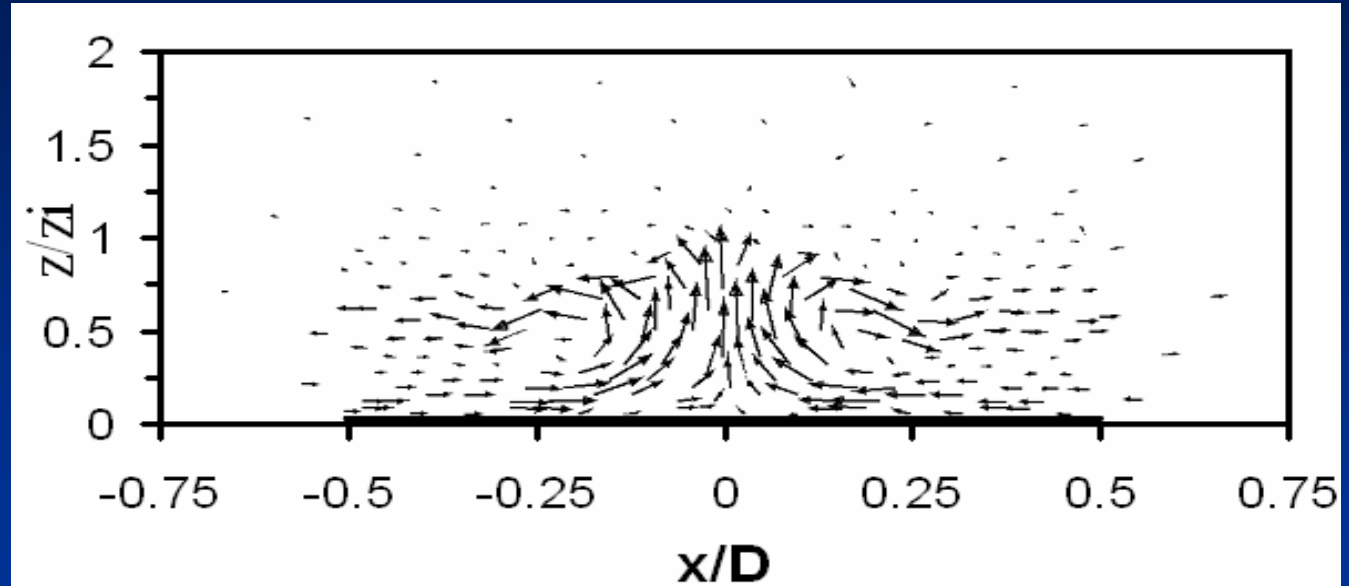


Thermal convection (UHIC)

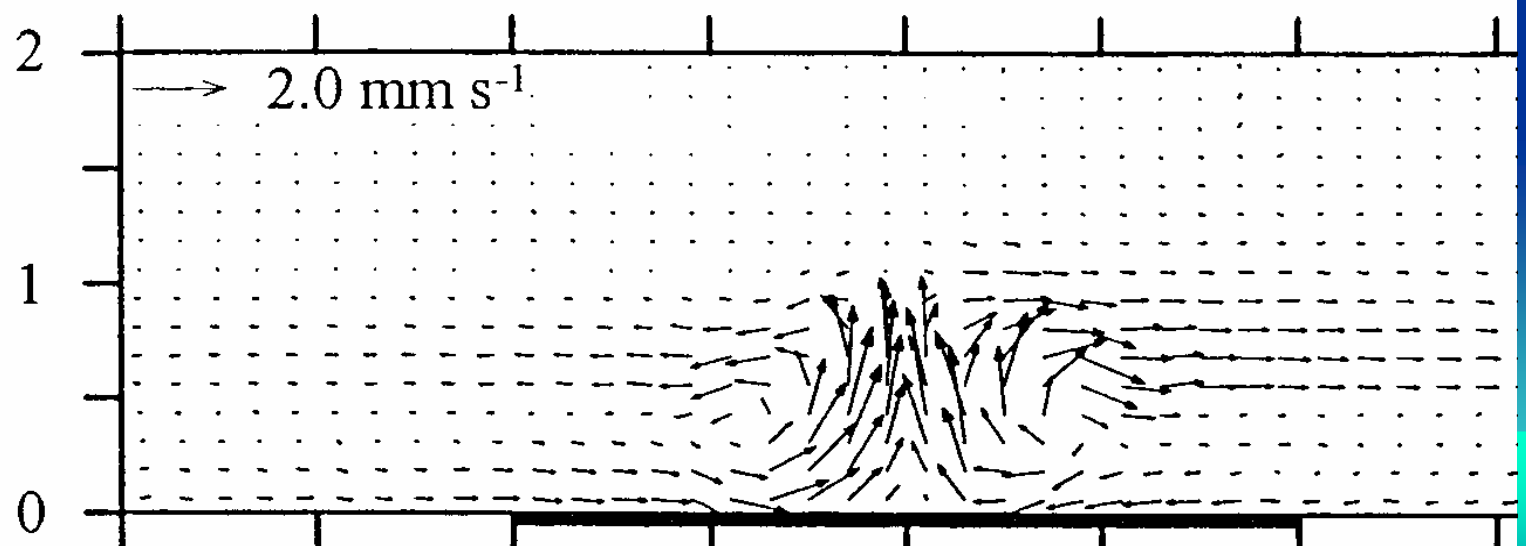
A.Cenedese, P.Monti: Interaction between an Inland Urban Heat Island and a Sea-Breeze Flow: A Laboratory Study, 2003.

CFD results

T(z) profiles are also in line with the measured data.



PIV results
(Cenedese & Monti 2003)



Down-burst test case

Straka et al.1990, Reinert 2007 

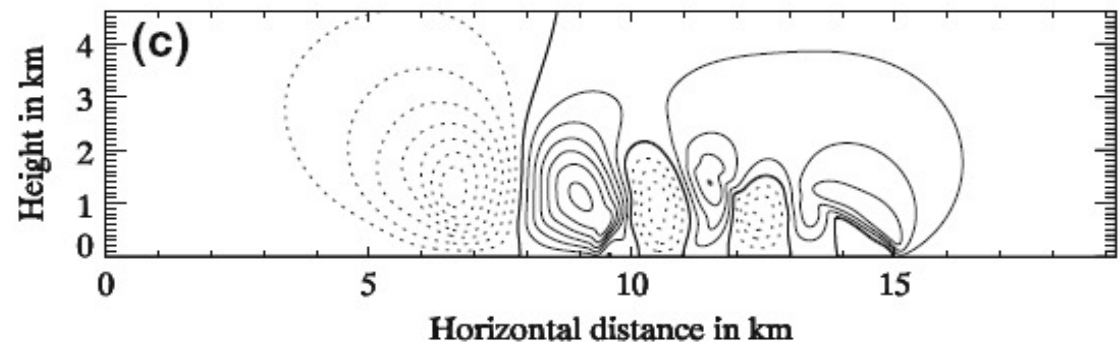
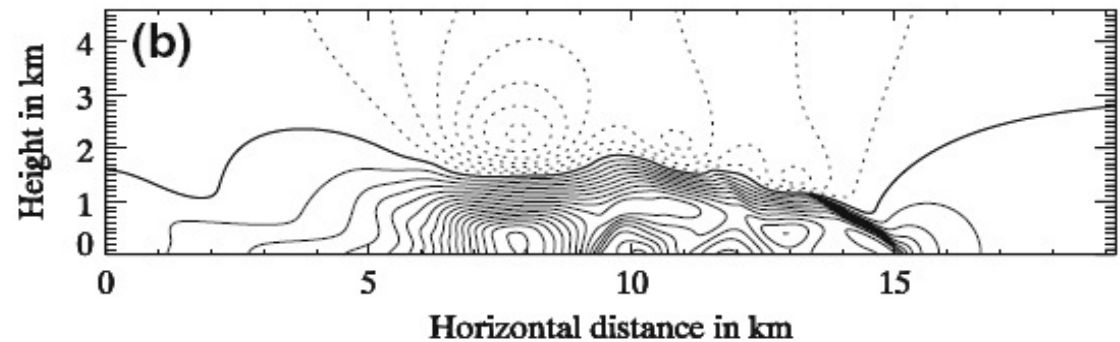
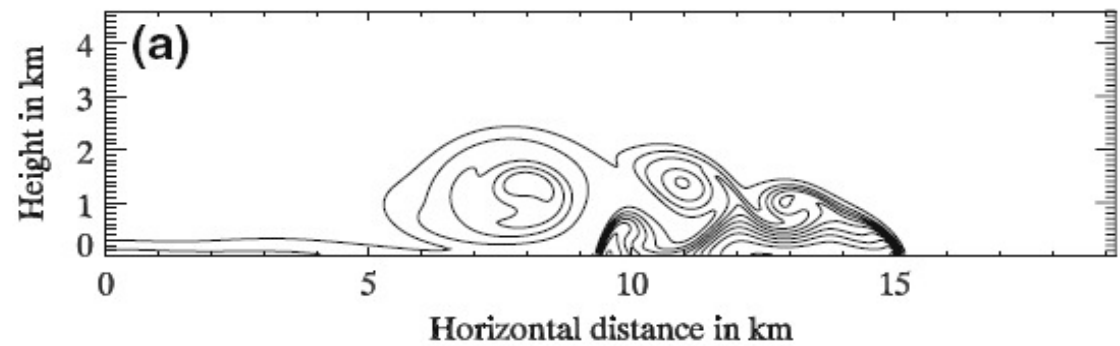
Cold bubble: -15°C



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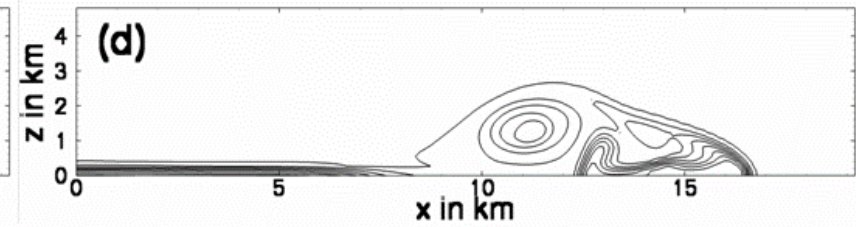
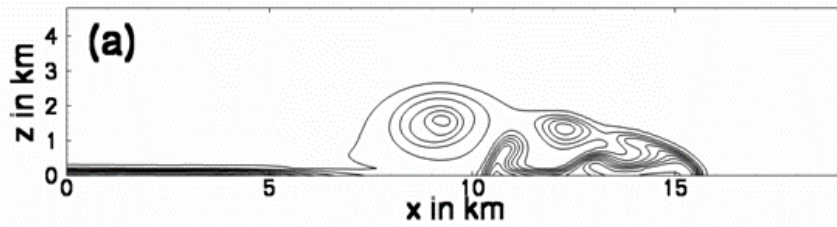


Results

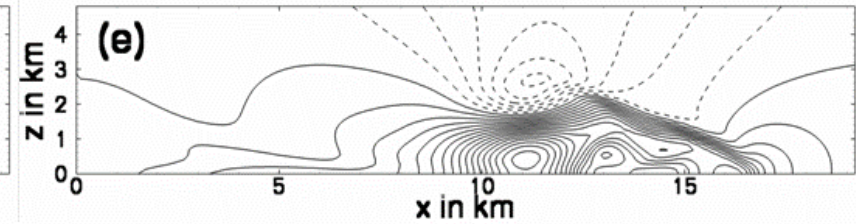
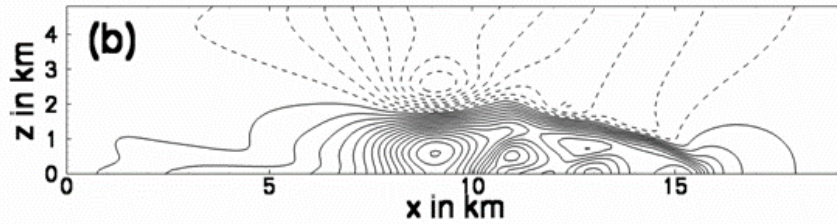
Compressible version

Simplified (incompressible)

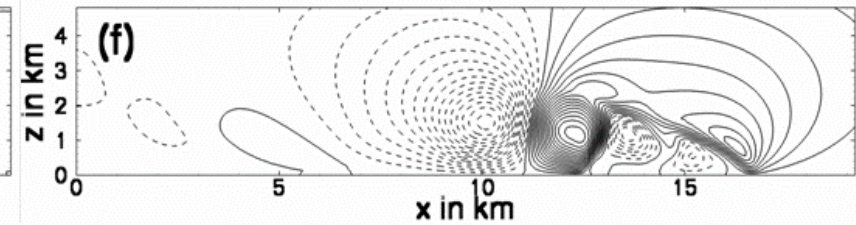
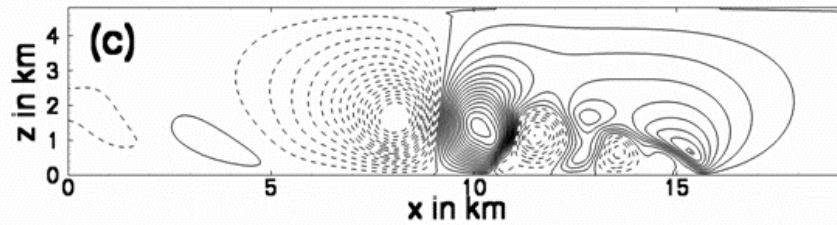
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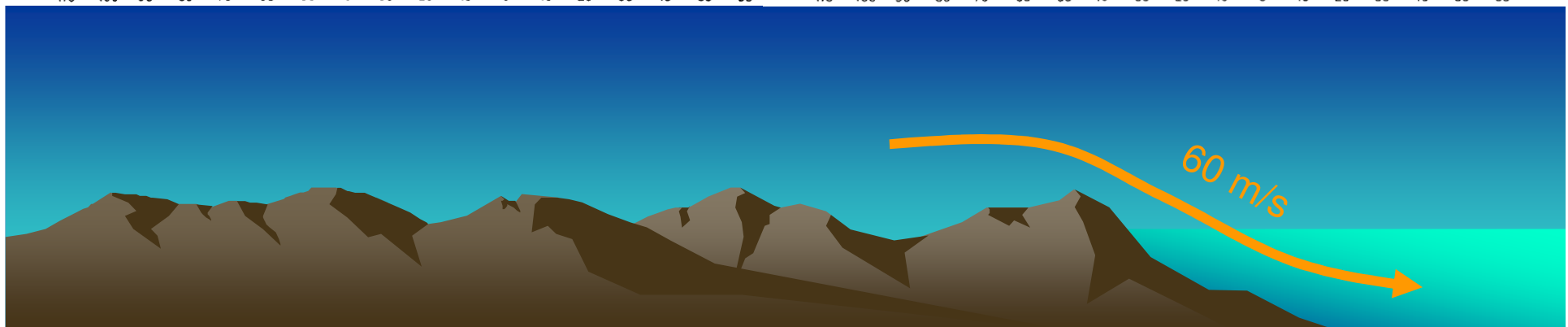
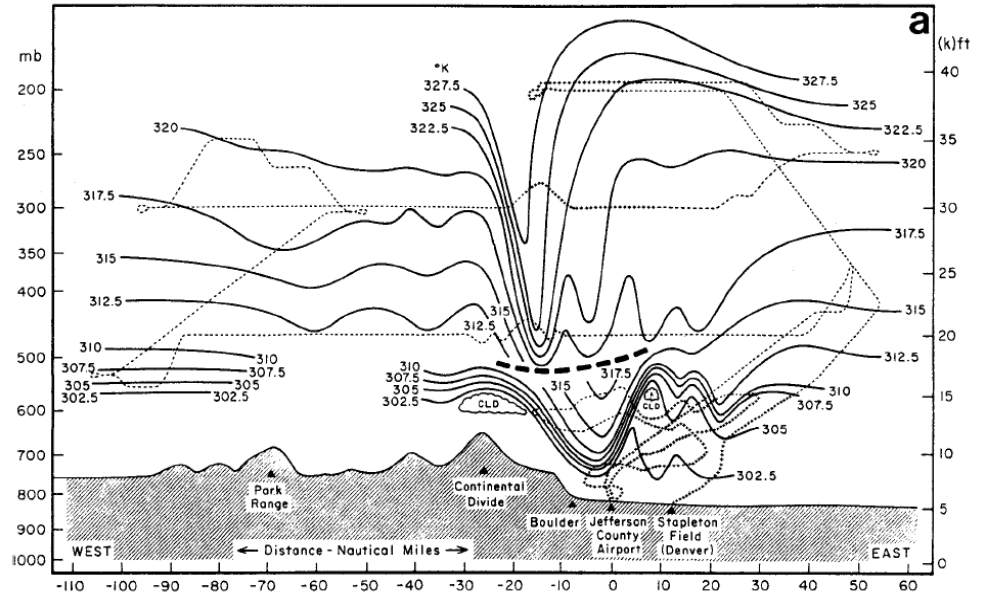
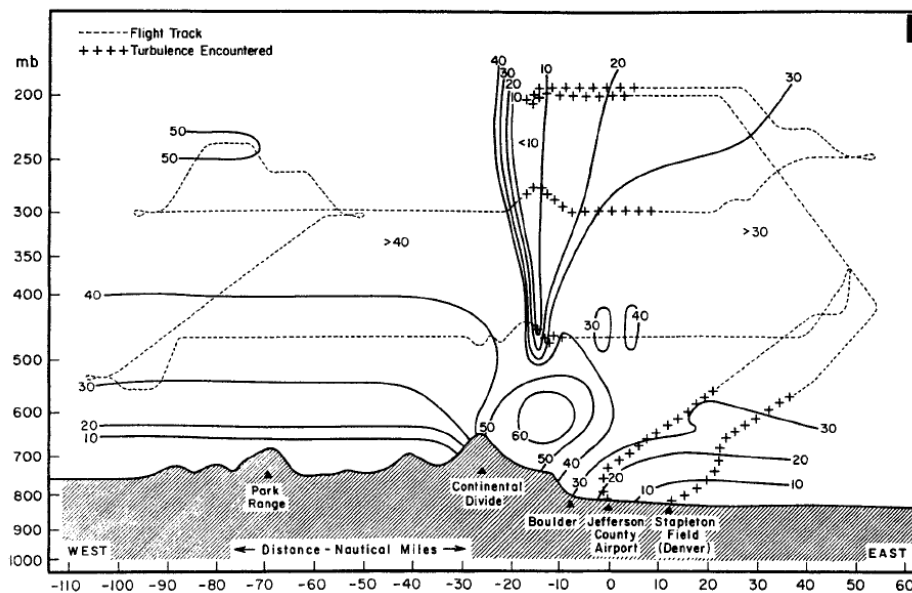


Down-slope windstorm

Boulder 1972 jan.

Measured velocity field

Measured potential temperature

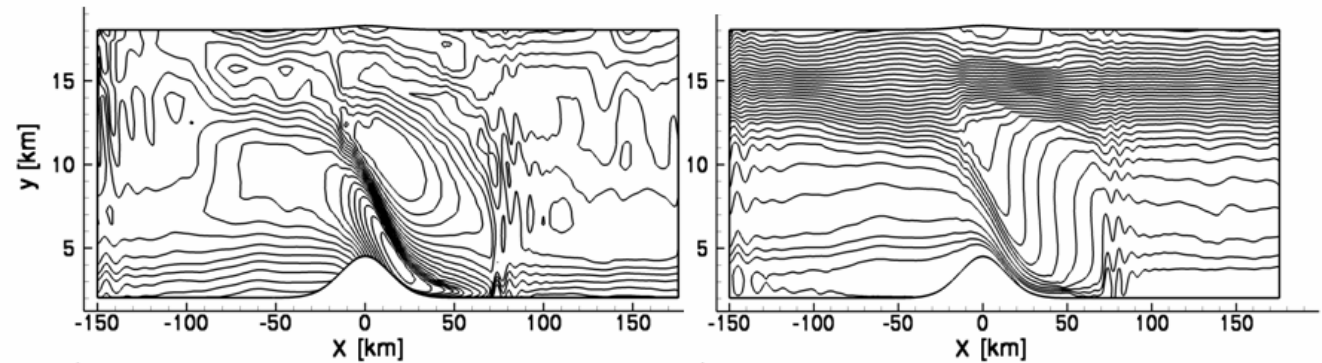


Results

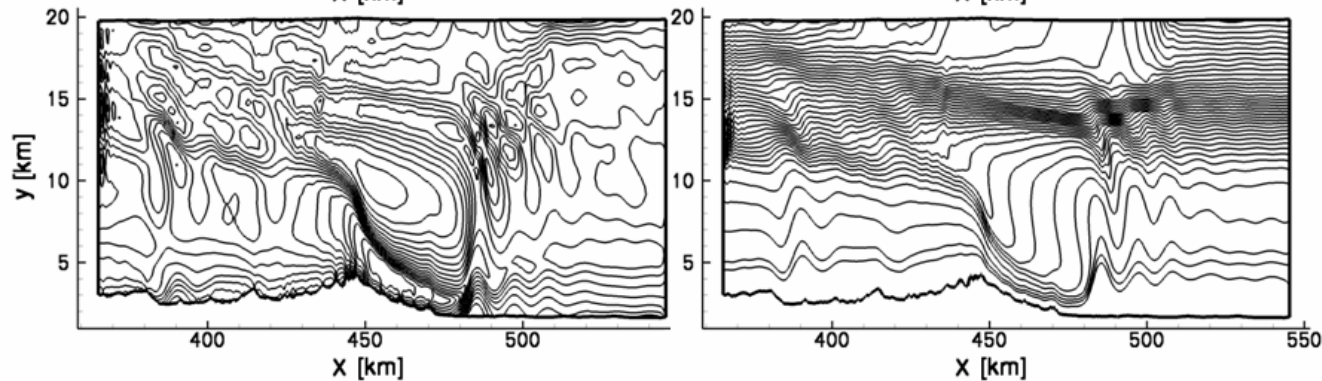
Velocity field

Potential temperature

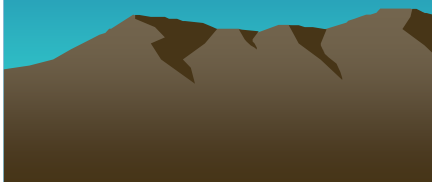
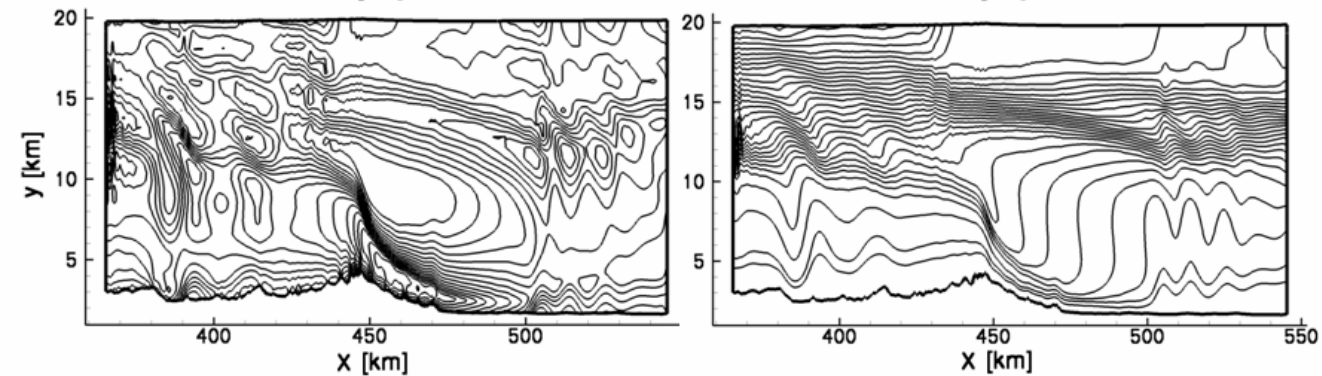
For idealized topography:



For realistic topography:



After a longer period of time:

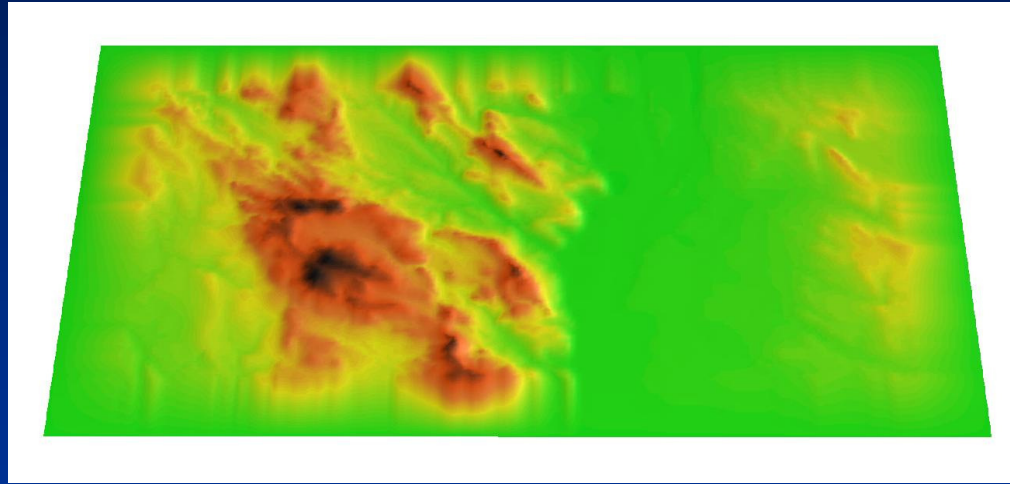


Two application examples

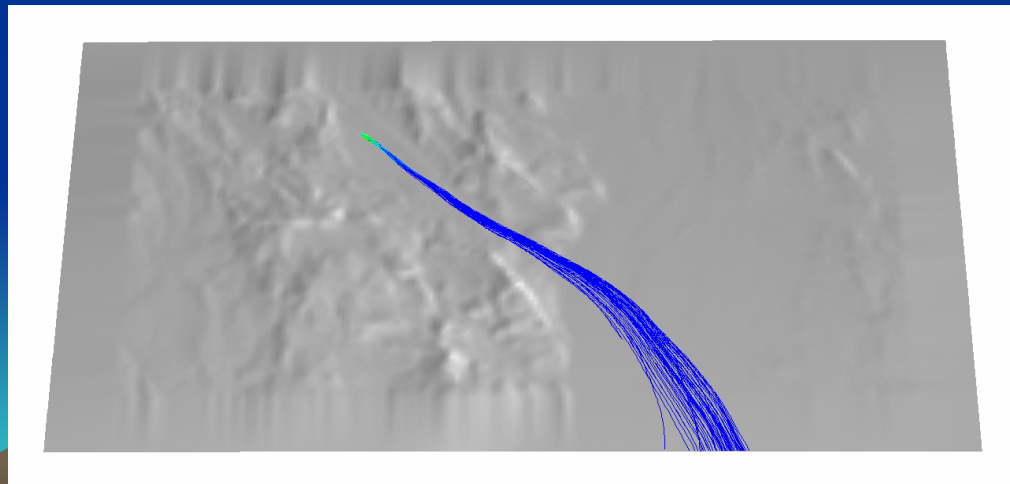
- Dispersion of pollutants
- Analyses of instabilities



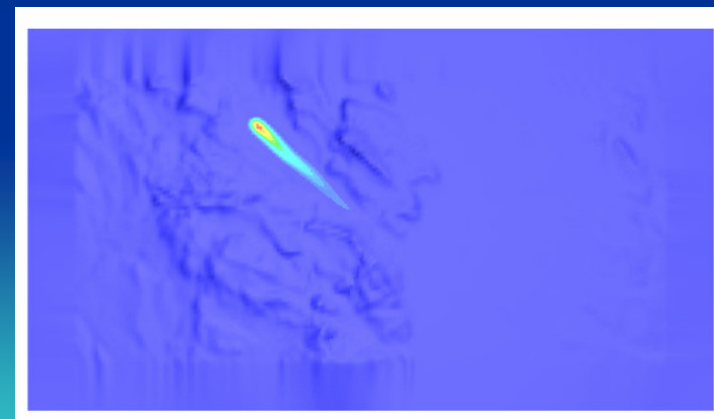
Meso scale atmospheric dispersion



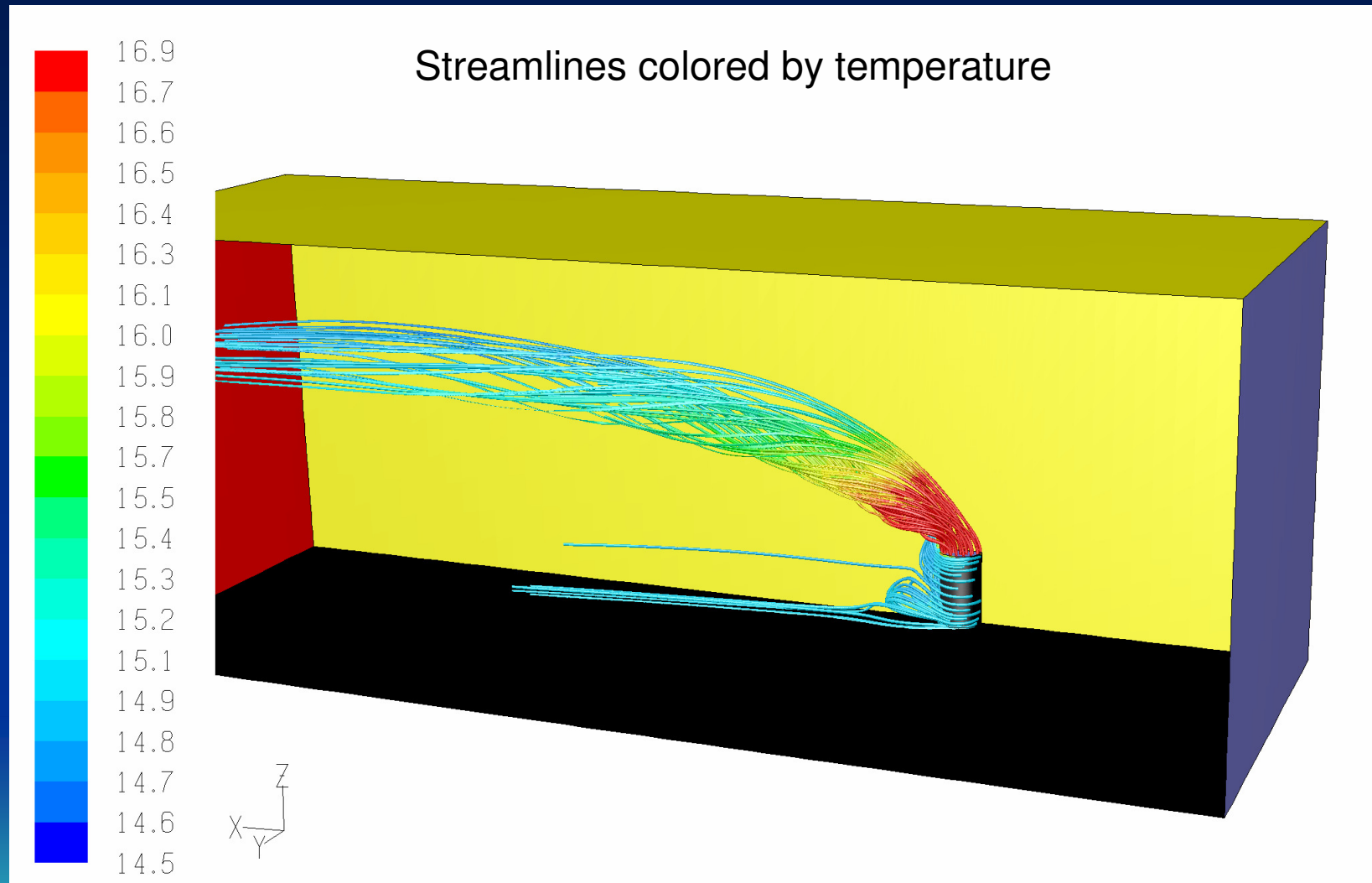
Orography of Pilis mountain



Evolution of surface concentration



Micro-scale atmospheric dispersion



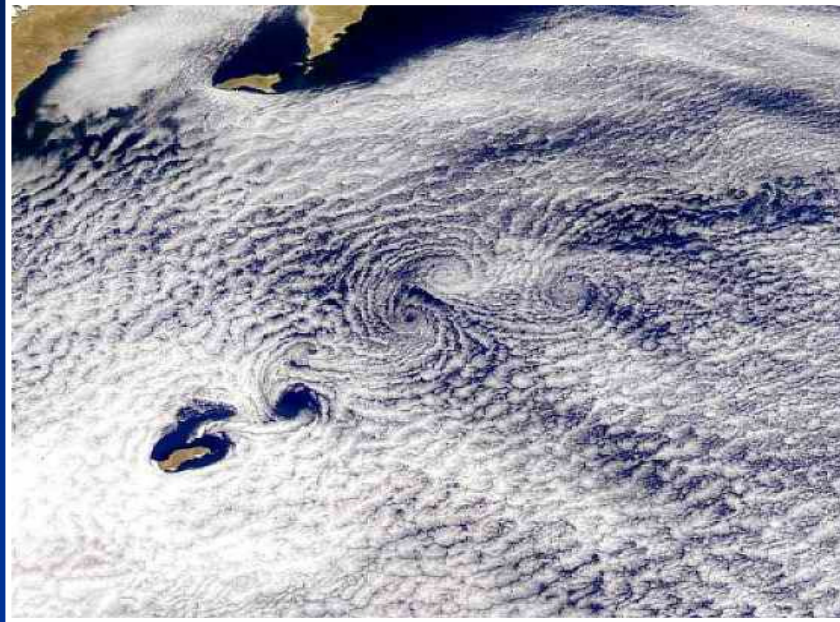
Chimney height 180 m

Standard (stable) temperature profile

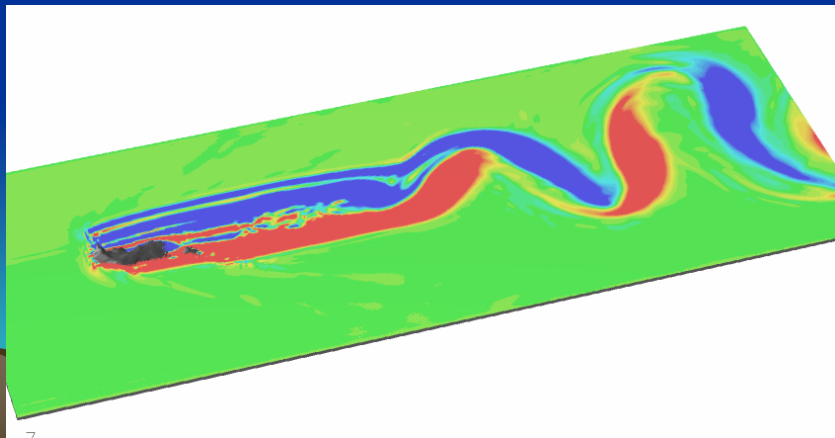
Wind speed: 3m/s

Injection velocity: 5 m/s

Von Kármán vortices behind a volcanic island



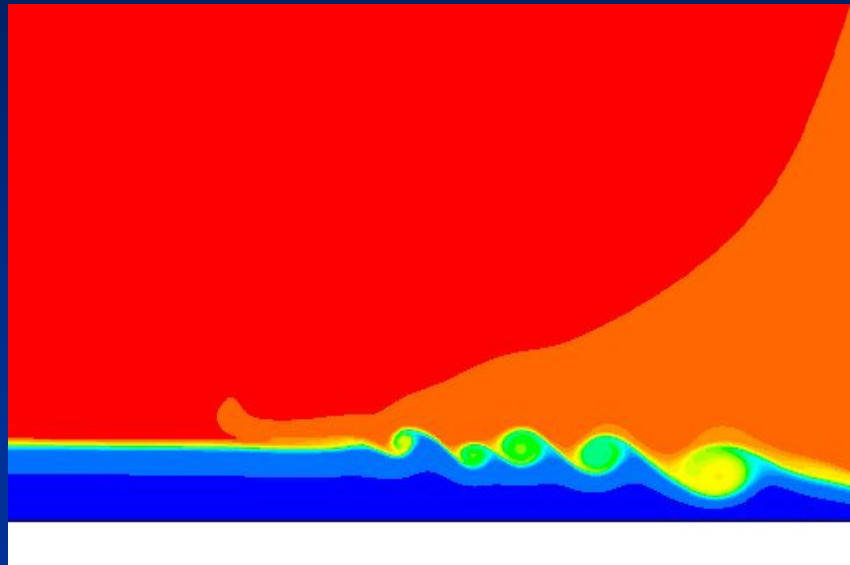
Satellite image about
Guadalupe island



First CFD results

Investigation of instabilities

Kelvin-Helmholtz instability



Comp. domain: 25 km x 5.5 km
Temperature difference 20 °C

Cloud formation:



Conclusions

- An easy to implement method has been developed for taking into account:
 - stratification effects,
 - adiabatic heat,
 - Coriolis force,
 - compressibility.
- The model has been validated against:
 - some analytic solutions,
 - laboratory experiments,
 - reference calculations,
 - in field measurements.
- Further effort is necessary for including:
 - moisture transport and phase changes,
 - porous drag models,
 - radiation heat transfer,
 - surface energy balance.
- Foreseeable applications:
 - local convections (e.g. UHIC, see breeze, valley breeze),
 - dispersion of pollutants (e.g. due to traffic, industry, chemical vapors),
 - meteorological research (e.g. gravity waves, cloud formation),
 - assessment of the wind power potential,
 - simulation of catastrophes (e.g. large fires, volcanism).

